

Low-Dimensional Reduced-Order Models for Statistical Response and Uncertainty Quantification in Turbulent Systems

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joint work with Di Qi

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Challenges for turbulent dynamical systems

Uncertainty quantification (UQ) deals with the **probabilistic characterization** of all the possible evolutions of a dynamical system given an *initial set of possible states* as well as *the random forcing or parameters*.

- Turbulent dynamical systems are characterized by *a large dimensional phase space* and *high degrees of internal instability*.
- Instabilities through *energy-conserving nonlinear interactions* result in a statistical steady state that is usually *non-Gaussian*.
- Accurate quantification for the statistical variability to *general external perturbations* is important in climate change sciences.

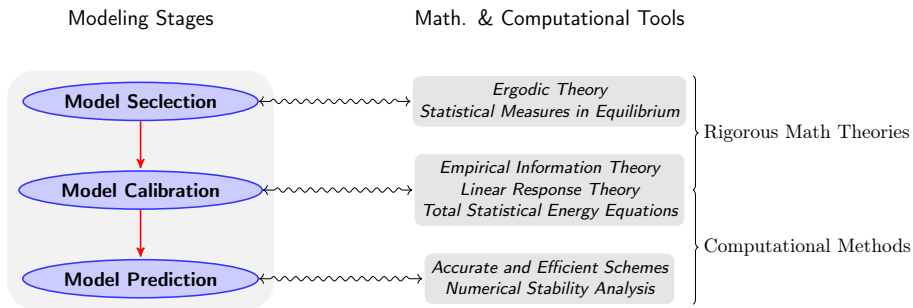
Major Task of this work:

Investigate a concise systematic framework for **measuring** and **optimizing consistency** and **sensitivity** of imperfect dynamical models.

General framework for statistical modeling

The system setup will be a finite-dimensional system of, $\mathbf{u} \in \mathbb{R}^N$, with **linear dynamics** and an **energy preserving quadratic part**

$$\frac{d\mathbf{u}}{dt} = \mathcal{L}[\mathbf{u}] = (L + D)\mathbf{u} + B(\mathbf{u}, \mathbf{u}) + \mathbf{F}(t) + \sigma_k(t)\dot{W}_k(t; \omega), \quad (1)$$
$$L^* = -L; \quad D \leq 0; \quad \mathbf{u} \cdot B(\mathbf{u}, \mathbf{u}) \equiv 0.$$



Related Works and Papers

- Recent new developments

- ▶ Majda, *Introduction to turbulent dynamical systems in complex systems*, Springer, 2016.
- ▶ Majda and Qi, *New strategies for reduced-order models for predicting the statistical responses and uncertainty quantification in complex turbulent dynamical systems*, SIAM Review, 2017.

- Statistical theories

- ▶ Majda, *Statistical energy conservation principle for inhomogeneous turbulent dynamical systems*. PNAS, 2015.
- ▶ Majda and Gershgorin, *Link between statistical equilibrium fidelity and forecasting skill for complex systems with model error*. PNAS, 2011.
- ▶ Majda and Wang, *Linear response theory for statistical ensembles in complex systems with time-periodic forcing*. CMS, 2010.

- Improving imperfect model skill

- ▶ Majda and Qi, *Improving prediction skill of imperfect turbulent models through statistical response and information theory*, Journal of Nonlinear Science, 2015.
- ▶ Qi and Majda, *Low-dimensional reduced-order models for statistical response and uncertainty quantification: two-layer baroclinic turbulence*, JAS, 2016.
- ▶ Qi and Majda, *Low-dimensional reduced-order models for statistical response and uncertainty quantification: barotropic turbulence with topography*, Physica D, 2016.

Outline

- 1 A two-layer quasi-geostrophic model for baroclinic turbulence
 - Two-layer baroclinic turbulence in ocean and atmosphere regimes
- 2 Reduced-order statistical models for general turbulent systems
 - Formulation of the exact statistical moment dynamics
 - A reduced-order statistical model with consistency and sensitivity
 - Model calibration for optimal performance
- 3 Low-dimensional reduced-order models for the two-layer system
 - Statistical equations for the two-layer model
 - Tuning imperfect model parameters in the training phase
 - Imperfect model prediction in various dynamical regimes

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The two-layer flow with forcing and dissipation

The **two-layer quasi-geostrophic model** with baroclinic instability is one simple but fully nonlinear fluid model capable in capturing the essential physics *in ocean and atmosphere science*.

Two-layer model

$$\begin{aligned}\frac{\partial q_\psi}{\partial t} + J(\psi, q_\psi) + J(\tau, q_\tau) + \beta \frac{\partial \psi}{\partial x} + U \frac{\partial \Delta \tau}{\partial x} &= -\frac{\kappa}{2} \Delta (\psi - \tau) - \nu \Delta^s q_\psi + \mathcal{F}_\psi, \\ \frac{\partial q_\tau}{\partial t} + J(\psi, q_\tau) + J(\tau, q_\psi) + \beta \frac{\partial \tau}{\partial x} + U \frac{\partial}{\partial x} (\Delta \psi + k_d^2 \psi) &= -\frac{\kappa}{2} \Delta (\tau - \psi) - \nu \Delta^s q_\tau + \mathcal{F}_\tau.\end{aligned}$$

Barotropic and baroclinic modes:

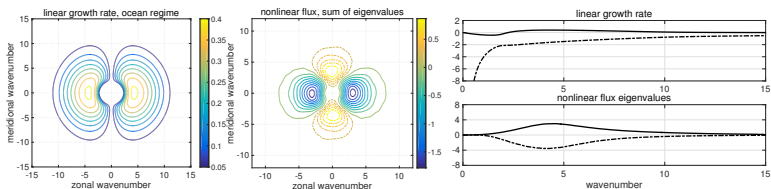
$$\begin{aligned}q_\psi &= \Delta \psi, \quad \psi = \frac{1}{2} (\psi_1 + \psi_2), \\ q_\tau &= \Delta \tau - k_d^2 \tau, \quad \tau = \frac{1}{2} (\psi_1 - \psi_2).\end{aligned}$$

Normalized energy-consistent modes:

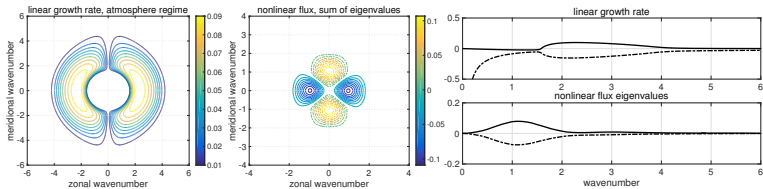
$$\begin{aligned}\rho_{\psi, \mathbf{k}} &= \frac{q_{\psi, \mathbf{k}}}{|\mathbf{k}|} = -|\mathbf{k}| \psi_{\mathbf{k}}, \\ \rho_{\tau, \mathbf{k}} &= \frac{q_{\tau, \mathbf{k}}}{\sqrt{|\mathbf{k}|^2 + k_d^2}} = -\sqrt{|\mathbf{k}|^2 + k_d^2} \tau_{\mathbf{k}}.\end{aligned}$$

Flow in high-latitude homogeneous regimes

regime	N	β	k_d	U	κ	(k_{\min}, k_{\max})	σ_{\max}	$(k_x, k_y)_{\max}$
ocean, high lat.	256	10	10	1	9	(2.25, 14.61)	0.411	(4, 0)
atmosphere, high lat.	256	1	4	0.2	0.2	(1.58, 6.78)	0.099	(2, 0)

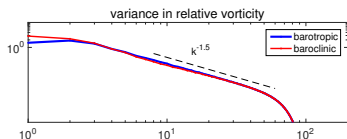
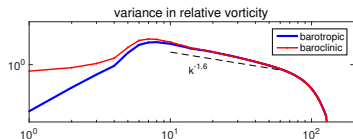
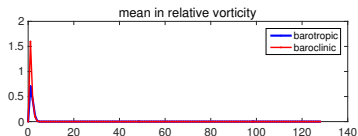
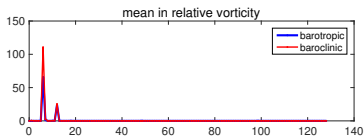
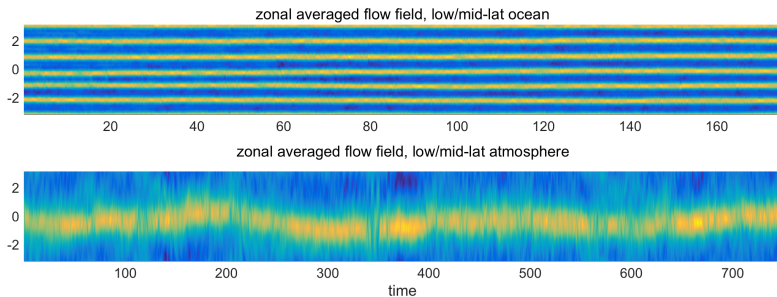


(a) high-latitude ocean regime



(b) high-latitude atmosphere regime

Flow in low/mid-latitude regimes with zonal jets



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General setup of turbulent systems with quadratic nonlinearities

The system setup will be a finite-dimensional system with linear dynamics and an energy preserving quadratic part with $\mathbf{u} \in \mathbb{R}^N$

$$\frac{d\mathbf{u}}{dt} = \mathcal{L}[\mathbf{u}(t; \omega); \omega] = (L + D)\mathbf{u} + B(\mathbf{u}, \mathbf{u}) + \mathbf{F}(t) + \sigma_k(t)\dot{W}_k(t; \omega), \quad (2)$$

$$\mathbf{u}(t_0; \omega) = \mathbf{u}_0(\omega). \quad (3)$$

- L being a **skew-symmetric linear operator** $L^* = -L$, representing the β -effect of Earth's curvature, topography etc.
- D being a **negative definite symmetric operator** $D^* = D$, representing dissipative processes such as surface drag, radiative damping, viscosity etc.
- $B(\mathbf{u}, \mathbf{u})$ being a **quadratic operator** which conserves the energy by itself so that it satisfies $B(\mathbf{u}, \mathbf{u}) \cdot \mathbf{u} = 0$.
- $\mathbf{F}(t) + \sigma_k(t)\dot{W}_k(t; \omega)$ being the **effects of external forcing**, i.e. solar forcing, seasonal cycle, which can be split into a mean component $\mathbf{F}(t)$ and a stochastic component with white noise characteristics.

Exact statistical moment equations

Statistical mean and covariance dynamics, $\mathbf{u} = \bar{\mathbf{u}} + Z_i \mathbf{v}_i$, $R_{ij} = \langle Z_i Z_j^* \rangle$,

$$\begin{aligned}\frac{d\bar{\mathbf{u}}}{dt} &= (L + D)\bar{\mathbf{u}} + B(\bar{\mathbf{u}}, \bar{\mathbf{u}}) + R_{ij} B(\mathbf{v}_i, \mathbf{v}_j) + \mathbf{F}(t), \\ \frac{dR}{dt} &= L_v R + R L_v^* + Q_F + Q_\sigma.\end{aligned}$$

- the linear dynamics operator L_v expressing energy transfers between the *mean field* and the *stochastic modes* (B), as well as *energy dissipation* (D), and *non-normal dynamics* (L)

$$\{L_v\}_{ij} = [(L + D)\mathbf{v}_j + B(\bar{\mathbf{u}}, \mathbf{v}_j) + B(\mathbf{v}_j, \bar{\mathbf{u}})] \cdot \mathbf{v}_i.$$

- the positive definite operator Q_σ expressing energy transfer due to external stochastic forcing

$$\{Q_\sigma\}_{ij} = \mathbf{v}_i^* \sigma_k^* \sigma_k \mathbf{v}_j.$$

- the third-order moments expressing the energy flux between different modes due to non-linear terms

$$Q_F = \langle Z_m Z_n Z_j \rangle B(\mathbf{v}_m, \mathbf{v}_n) \cdot \mathbf{v}_i + \langle Z_m Z_n Z_i \rangle B(\mathbf{v}_m, \mathbf{v}_n) \cdot \mathbf{v}_j.$$

note that energy is still conserved in this nonlinear interaction part

$$\begin{aligned}\text{Tr}[Q_F] &= 2 \langle Z_m Z_n Z_i \rangle B(\mathbf{v}_m, \mathbf{v}_n) \cdot \mathbf{v}_i \\ &= 2 \langle B(Z_m \mathbf{v}_m, Z_n \mathbf{v}_n) \cdot Z_i \mathbf{v}_i \rangle = 2 \langle B(\mathbf{u}', \mathbf{u}') \cdot \mathbf{u}' \rangle = 0.\end{aligned}$$

Reduced-Order Statistical Energy Closure

The true statistical model

$$\frac{d\bar{\mathbf{u}}}{dt} = (L + D)\bar{\mathbf{u}} + B(\bar{\mathbf{u}}, \bar{\mathbf{u}}) + R_{ij}B(\mathbf{v}_i, \mathbf{v}_j) + \mathbf{F}(t), \quad \bar{\mathbf{u}} \in \mathbb{R}^N,$$

$$\frac{dR}{dt} = L_V(\bar{\mathbf{u}})R + RL_V^*(\bar{\mathbf{u}}) + Q_F + Q_\sigma, \quad R \in \mathbb{R}^{N \times N}.$$

$$Q_{F,ij} = \langle Z_m Z_n Z_j \rangle B(\mathbf{v}_m, \mathbf{v}_n) \cdot \mathbf{v}_i + \langle Z_m Z_n Z_i \rangle B(\mathbf{v}_m, \mathbf{v}_n) \cdot \mathbf{v}_j$$

Reduced-Order Statistical Energy Closure

The reduced-order approximation $\bar{\mathbf{u}}_M \in \mathbb{R}^M$, $M \ll N$

$$\begin{aligned}\frac{d\bar{\mathbf{u}}_M}{dt} &= (L+D)\bar{\mathbf{u}}_M + B(\bar{\mathbf{u}}_M, \bar{\mathbf{u}}_M) + R_{M,ij}B(\mathbf{v}_i, \mathbf{v}_j) + \mathbf{F}, \\ \frac{dR_M}{dt} &= L_V R_M + R_M L_V^* + Q_F^M + Q_\sigma,\end{aligned}$$

A preferred approach for the nonlinear flux Q_F^M combining both the *detailed model energy mechanism* and *control over model sensitivity* is proposed

$$Q_F^M = Q_F^{M,-} + Q_F^{M,+} = f_1(E) [-(N_{M,\text{eq}} + d_M l_N) R_M] + f_2(E) [Q_{F,\text{eq}}^+ + \Sigma_M].$$

Reduced-Order Statistical Energy Closure

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- Higher-order corrections from equilibrium statistics:

$$Q_{F,\text{eq}} = Q_{F,\text{eq}}^- + Q_{F,\text{eq}}^+ = -L_V(\bar{\mathbf{u}}_{\text{eq}}) R_{\text{eq}} - R_{\text{eq}} L_V^*(\bar{\mathbf{u}}_{\text{eq}}) - Q_\sigma, \quad N_{M,\text{eq}} = \frac{1}{2} Q_{F,\text{eq}}^- R_{\text{eq}}^{-1}.$$

- Additional damping and noise to model nonlinear flux:

$$Q_M^{\text{add}} = -d_M R_M + \Sigma_M.$$

- Statistical energy-consistent scaling to improve model sensitivity:

$$f_1(E) = \left(\frac{E}{E_{\text{eq}}} \right)^{1/2}, \quad f_2(E) = \left(\frac{E}{E_{\text{eq}}} \right)^{3/2}.$$

Climate fidelity for equilibrium

Equilibrium fidelity refers to the convergence to the same final unperturbed statistical equilibrium R_{eq} in the reduced-order models R_M in each resolved component.

Specifically, it requires that the model nonlinear flux correction term Q_M converges to the truth, $Q_M \rightarrow Q_{F,\text{eq}}$, when no external perturbation is added

$$\frac{dR_{M,\text{eq}}}{dt} = 0 = L_V(\bar{\mathbf{u}}_{\text{eq}})R_{M,\text{eq}} + R_{M,\text{eq}}L_V^*(\bar{\mathbf{u}}_{\text{eq}}) + Q_{F,\text{eq}}^M + Q_\sigma \rightarrow R_{M,\text{eq}} = R_{\text{eq}}.$$

- the first component $(N_{M,\text{eq}}, Q_{F,\text{eq}}^+)$ comes from the true equilibrium statistics.
- climate consistency requires the second component correction makes no contribution in the unperturbed case

$$\Sigma_M = \frac{1}{2}d_M R_{\text{eq}}, \quad f_1(E_{\text{eq}}) = 1, \quad f_2(E_{\text{eq}}) = 1.$$

Statistical energy conservation principle¹

Theorem

(Statistical Energy Conservation Principle) Under the structural symmetries on the basis \mathbf{v}_i , for any turbulent dynamical systems in the form (1) the total statistical energy, $E = \bar{E} + E' = \frac{1}{2} \bar{\mathbf{u}} \cdot \bar{\mathbf{u}} + \frac{1}{2} \text{tr}R$, satisfies

$$\frac{dE}{dt} = \bar{\mathbf{u}} \cdot D\bar{\mathbf{u}} + \bar{\mathbf{u}} \cdot \mathbf{F} + \text{tr}(DR) + \frac{1}{2} \text{tr}Q_\sigma,$$

where R satisfies the exact covariance equation.

Corollary

Under the assumption of the Theorem, assume $D = -dI$, with $d > 0$, then the turbulent dynamical system satisfies the closed statistical energy equation for $E = \frac{1}{2} \bar{\mathbf{u}} \cdot \bar{\mathbf{u}} + \frac{1}{2} \text{tr}R$,

$$\frac{dE}{dt} = -2dE + \bar{\mathbf{u}} \cdot \mathbf{F} + \frac{1}{2} \text{tr}Q_\sigma.$$

¹Majda, *Statistical energy conservation principle for inhomogeneous turbulent dynamical systems*, PNAS, 2015.

Model calibration blending *statistical response* and *information theory*

Accurate modeling about the model sensitivity to various external perturbations requires the imperfect reduced-order models to correctly reflect the true system's “memory” about its previous states.

- *the linear response operator* can characterize the model sensitivity involving the nonlinear effects in the system regardless of the specific forms of the external perturbations.
- *empirical information theory* can be used as the distance between these two operators to calculate the unbiased and invariant measure for model distributions.

Linear response operator $\mathcal{R}_A(t)$

Linear response theory: The linear response of a system, $\mathbf{u}_t = \mathbf{f}(\mathbf{u})$, $\mathbf{f}^\delta = \mathbf{f} + \delta \mathbf{f}'(t)$, can be predicted by observing appropriate statistics of the system in **equilibrium** π_{eq}

$$\begin{aligned}\mathbb{E}^\delta A(\mathbf{u}) &= E_{\text{eq}}(A) + \delta E'_A + O(\delta^2), \\ \delta E'_A &= \int_0^t \mathcal{R}_A(t-s) \delta f'(s) ds.\end{aligned}$$

without the need of applying any perturbations.

kicked response: For δ small enough, the linear response operator $\mathcal{R}_A(t)$ can be calculated by solving the **unperturbed system** with a perturbed initial distribution

$$\pi|_{t=0} = \pi_{\text{eq}}(\mathbf{u} - \delta \mathbf{u}) = \pi_{\text{eq}} - \delta \mathbf{u} \cdot \nabla \pi_{\text{eq}} + O(\delta^2).$$

$$\delta \mathcal{R}_A(t) \equiv \delta \mathbf{u} \cdot \mathcal{R}_A = \int A(\mathbf{u}) \delta \pi' + O(\delta^2).$$

Link between equilibrium fidelity and forecasting skill

Given the optimal model for the unperturbed climate π_{eq} , how can we assess the error in the climate change prediction

$$\mathcal{P}(\pi, \pi^M) = \int \pi \ln \frac{\pi}{\pi^M},$$

based on the unperturbed climate?

Under assumptions with diagonal covariance matrices $R = \text{diag}(R_k)$ and equilibrium model fidelity $\mathcal{P}(\pi_G, \pi_G^M) = 0$

$$\begin{aligned} \mathcal{P}(\pi_\delta, \pi_\delta^M) &= \mathcal{S}(\pi_{G,\delta}) - \mathcal{S}(\pi_\delta) \\ &+ \frac{1}{2} \sum_k (\delta \bar{u}_k - \delta \bar{u}_{M,k}) R_k^{-1} (\delta \bar{u}_k - \delta \bar{u}_{M,k}) \\ &+ \frac{1}{4} \sum_k R_k^{-2} (\delta R_k - \delta R_{M,k})^2 + O(\delta^3). \end{aligned}$$

R_k is the equilibrium variance in k -th component, and $\delta \bar{u}_k$ and δR_k are the linear response operators for the mean and variance in k -th component.

(Majda & Gershgorin, PNAS, 2011)

Optimization in the training phase

To summarize, consider a class of imperfect models, \mathcal{M} . The *optimal model* $M^* \in \mathcal{M}$ that ensures best information consistent responses is characterized with the smallest additional information in the linear response operator \mathcal{R}_A , such that

$$\left\| \mathcal{P} \left(p_\delta, p_\delta^{M^*} \right) \right\|_{L^1([0, T])} = \min_{M \in \mathcal{M}} \left\| \mathcal{P} \left(p_\delta, p_\delta^M \right) \right\|_{L^1([0, T])},$$

- p_δ^M can be achieved through a kicked response procedure in the training phase compared with the actual observed data p_δ in nature;
- the information distance between perturbed responses $\mathcal{P} \left(p_\delta, p_\delta^M \right)$ can be calculated through the expansion formula;
- the information distance $\mathcal{P} \left(p_\delta(t), p_\delta^M(t) \right)$ is measured at each time instant, so the entire error is averaged under the L^1 -norm inside a time window $[0, T]$.

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Exact statistical moment equations for the two-layer model

The rescaled set of equations of (1) can be summarized in the abstract form

$$\frac{d\mathbf{p}_k}{dt} = B_k(\mathbf{p}_k, \mathbf{p}_k) + (\mathcal{L}_k - \mathcal{D}_k)\mathbf{p}_k + \mathcal{F}_k, \quad \mathbf{p}_k = (\rho_{\psi, k}, \rho_{\tau, k})^T, \quad \sum_{\mathbf{k}} \mathbf{p}_k \cdot B_k(\mathbf{p}_k, \mathbf{p}_k) \equiv 0,$$

where the normalized state variable $\mathbf{p}_k = (\rho_{\psi, k}, \rho_{\tau, k})^T$ is in barotropic and baroclinic mode, the linear operator is decomposed into non-symmetric part \mathcal{L}_k involving β -effect and shear flow U and dissipation part \mathcal{D}_k , together with the forcing \mathcal{F}_k combining deterministic component and stochastic component.

$$\mathcal{L}_k = \begin{bmatrix} \frac{ik_x \beta}{|\mathbf{k}|^2} & -\frac{ik_x U}{\sqrt{1+(k_d/|\mathbf{k}|)^2}} \\ -ik_x U \frac{1-(k_d/|\mathbf{k}|)^2}{\sqrt{1+(k_d/|\mathbf{k}|)^2}} & \frac{ik_x \beta}{|\mathbf{k}|^2 + k_d^2} \end{bmatrix}, \quad \mathcal{D}_k = \frac{\kappa}{2} \begin{bmatrix} -1 & \frac{1}{\sqrt{1+(k_d/|\mathbf{k}|)^2}} \\ \frac{1}{\sqrt{1+(k_d/|\mathbf{k}|)^2}} & -\frac{1}{1+(k_d/|\mathbf{k}|)^2} \end{bmatrix},$$

$$B_k(\mathbf{p}_k, \mathbf{p}_k) = \begin{bmatrix} B_{\psi, k} \\ B_{\tau, k} \end{bmatrix} = \begin{bmatrix} \sum_{\mathbf{m}+\mathbf{n}=\mathbf{k}} \frac{\mathbf{m}^\perp \cdot \mathbf{n}}{|\mathbf{k}|} \left(\frac{|\mathbf{n}|}{|\mathbf{m}|} \rho_{\psi, \mathbf{m}} \rho_{\psi, \mathbf{n}} + \sqrt{\frac{|\mathbf{n}|^2 + k_d^2}{|\mathbf{m}|^2 + k_d^2}} \rho_{\tau, \mathbf{m}} \rho_{\tau, \mathbf{n}} \right) \\ \sum_{\mathbf{m}+\mathbf{n}=\mathbf{k}} \frac{\mathbf{m}^\perp \cdot \mathbf{n}}{\sqrt{|\mathbf{k}|^2 + k_d^2}} \left(\frac{\sqrt{|\mathbf{n}|^2 + k_d^2}}{|\mathbf{m}|} \rho_{\psi, \mathbf{m}} \rho_{\tau, \mathbf{n}} + \frac{|\mathbf{n}|}{\sqrt{|\mathbf{m}|^2 + k_d^2}} \rho_{\tau, \mathbf{m}} \rho_{\psi, \mathbf{n}} \right) \end{bmatrix}.$$

Exact statistical moment equations

Statistical energy in each spectral mode

$$R_{\mathbf{k}} = \overline{\mathbf{p}_{\mathbf{k}}^* \mathbf{p}_{\mathbf{k}}} = \begin{bmatrix} \overline{|p_{\psi, \mathbf{k}}|^2} & \overline{p_{\psi, \mathbf{k}}^* p_{\tau, \mathbf{k}}} \\ \overline{p_{\psi, \mathbf{k}} p_{\tau, \mathbf{k}}^*} & \overline{|p_{\tau, \mathbf{k}}|^2} \end{bmatrix}, \quad \overline{p_{1, \mathbf{k}}^* p_{2, \mathbf{k}}} = \bar{p}_{1, \mathbf{k}}^* \bar{p}_{2, \mathbf{k}} + \overline{p'_{1, \mathbf{k}} p'_{2, \mathbf{k}}}.$$

$R_{\mathbf{k}}$ combines the variability in both *mean* and *variance*. The true statistical dynamical equations form a 2×2 system about $R_{\mathbf{k}} \in \mathbb{C}^{2 \times 2}$

$$\frac{dR_{\mathbf{k}}}{dt} = (\mathcal{L}_{\mathbf{k}} - \mathcal{D}_{\mathbf{k}})R_{\mathbf{k}} + Q_{F, \mathbf{k}} + Q_{\sigma, \mathbf{k}} + \text{c.c.}, \quad |\mathbf{k}| \leq N,$$

$$Q_{F, \mathbf{k}} = \overline{\mathbf{p}_{\mathbf{k}}^* B_{\mathbf{k}}(\mathbf{p}_{\mathbf{k}}, \mathbf{p}_{\mathbf{k}})} = \begin{bmatrix} \overline{p_{\psi, \mathbf{k}}^* B_{\psi, \mathbf{k}}} & \overline{p_{\psi, \mathbf{k}}^* B_{\tau, \mathbf{k}}} \\ \overline{p_{\tau, \mathbf{k}}^* B_{\psi, \mathbf{k}}} & \overline{p_{\tau, \mathbf{k}}^* B_{\tau, \mathbf{k}}} \end{bmatrix}, \quad \sum_{\mathbf{k}} \text{tr} Q_{F, \mathbf{k}} \equiv 0.$$

Statistical energy conservation principle

The total statistical energy dynamical equation concerns the evolution of the *total variability in mean and variance* in response to external perturbations

$$E = \frac{1}{2} \sum_{1 \leq |\mathbf{k}| \leq N} |\mathbf{k}|^2 \overline{|\psi_{\mathbf{k}}|^2} + (|\mathbf{k}|^2 + k_d^2) \overline{|\tau_{\mathbf{k}}|^2} = \frac{1}{2} \sum_{1 \leq |\mathbf{k}| \leq N} \overline{|\rho_{\psi, \mathbf{k}}|^2} + \overline{|\rho_{\tau, \mathbf{k}}|^2}.$$

The exact dynamics for the statistical energy can be derived as

$$\frac{dE}{dt} + H_f = -\kappa E + \frac{\kappa}{2} F - \nu H + Q_{\sigma}.$$

H_f is the meridional heat flux due to baroclinic instability, F is the additional damping effects due to the non-symmetry in Ekman drag only applied on the bottom layer

$$H_f = k_d^2 U \int \overline{\psi_x \tau} = k_d^2 U \sum i k_x \overline{\psi_{\mathbf{k}}^* \tau_{\mathbf{k}}}, \quad F = \sum k_d^2 \overline{|\tau_{\mathbf{k}}|^2} + 2 \sum |\mathbf{k}|^2 \Re \overline{\psi_{\mathbf{k}}^* \tau_{\mathbf{k}}}.$$

Set-up for the numerical problem

- The true statistics is calculated by a pseudo-spectra code with 128 spectral modes zonally and meridionally, corresponding to $256 \times 256 \times 2$ grid points in total.
- In the reduced-order methods, only the large-scale modes $|\mathbf{k}| \leq 10$ are resolved, which is about 0.15% of the full model resolution.

External forcing in stochastic and deterministic component:

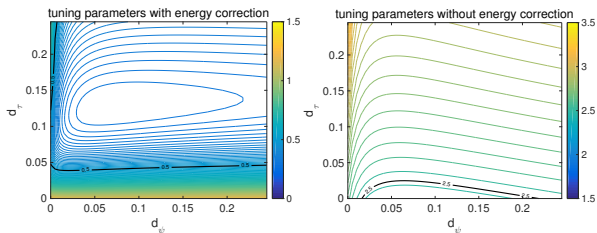
- The amplitude of the *stochastic forcing* $\sigma_k \dot{W}_k$ is introduced according to the equilibrium energy so that

$$\sigma_{\psi, \mathbf{k}}^2 = \delta \sigma_0^2 \overline{|q_{\psi, \mathbf{k}}|^2}_{\text{eq}}, \quad \sigma_{\tau, \mathbf{k}}^2 = \delta \sigma_0^2 \overline{|q_{\tau, \mathbf{k}}|^2}_{\text{eq}}.$$

- The *deterministic forcing* is introduced through a perturbation in the background shear $U_\delta = U + \delta U$

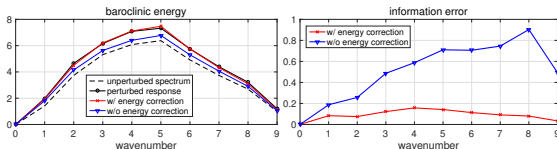
$$\delta f_{\psi, \mathbf{k}} = \delta U i k_x \left(-|\mathbf{k}|^2 \right) \tau_{\mathbf{k}}, \quad \delta f_{\tau, \mathbf{k}} = \delta U i k_x \left(-|\mathbf{k}|^2 + k_d^2 \right) \psi_{\mathbf{k}}.$$

Tuning parameters in the training phase



(a) tuning with energy scaling

(b) tuning without energy scaling

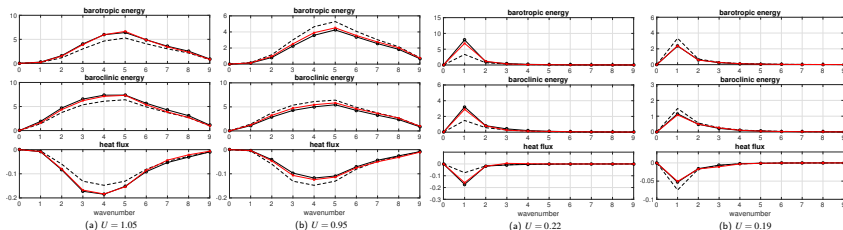


(c) prediction and info. errors

Figure: Tuning imperfect model parameters in the training phase. The information errors with varying model parameters, $d_M = (d_\psi, d_\tau)$, are plotted for stochastic barotropic perturbation case.

High-latitude: Mean shear flow perturbation

- The model is perturbed by changing the background zonal flow strength U ;
- The entire spectral is perturbed due to the mean flow advection in each spectral mode.

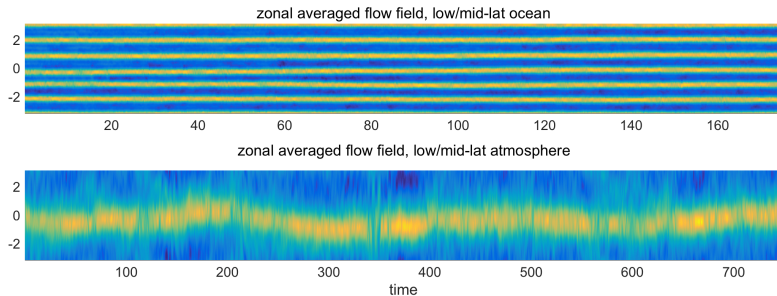


(a) ocean regime $\delta U = \pm 0.05$, $U_0 = 1$

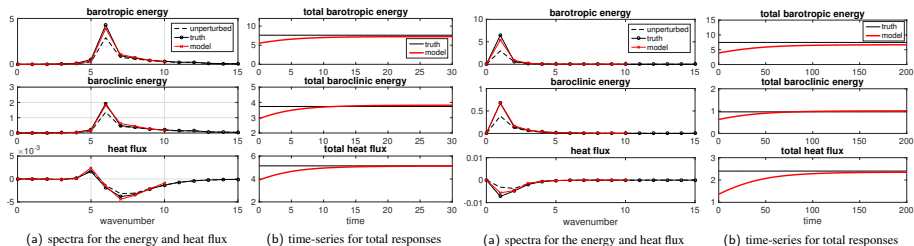
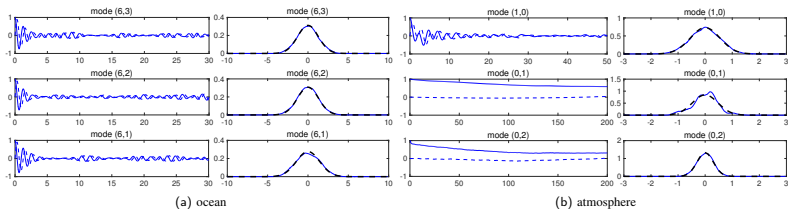
(b) atmos. regime $\delta U = \pm 0.02$, $U_0 = 0.2$

Flow in low-latitude regimes with zonal jets

regime	N	β	k_d	U	κ	(k_{\min}, k_{\max})	σ_{\max}	$(k_x, k_y)_{\max}$
ocean, high lat.	256	100	10	1	1	(7.14, 15.63)	0.104	(2, 8)
atmosphere, high lat.	256	2.5	4	0.2	0.05	(2.51, 7.06)	0.053	(3, 0)



Low-latitude: Stochastic perturbation with $\delta\sigma_0^2 = 0.2$ autocorrelation functions and probability density functions



Current and future work

- Reduced-order stochastic modeling strategies to capture passive scalar intermittency
 - ▶ Majda & Tong, *Intermittency in turbulent diffusion models with a mean gradient*, Nonlinearity, 2015
 - ▶ Majda & Gershgorin, *Elementary models for turbulent diffusion with complex physical features: eddy diffusivity, spectrum and intermittency*, Phil. Trans. R. Soc. A, 2013
 - ▶ Qi & Majda, *Predicting fat-tailed intermittent probability distributions in passive scalar turbulence with imperfect models through empirical information theory*, Comm. Math. Sci, 2015.
 - ▶ Qi & Majda, *Predicting Extreme Events for Passive Scalar Turbulence in Two-Layer Baroclinic Flows through Reduced-Order Stochastic Models*, Comm. Math. Sci, 2017.
- Design of a mitigation control strategy by using novel low-order statistical models;
 - ▶ Majda & Qi, *Effective control of complex turbulent dynamical systems through statistical functionals*, PNAS, 2017
 - ▶ Majda & Qi, *Using Statistical Functionals for Effective Control of Inhomogeneous Complex Turbulent Dynamical Systems*, submitted to Physica D.
- Rigorous statistical UQ for turbulent geophysical flows
 - ▶ Majda & Qi, Rigorous statistical uncertainty quantification for one-layer turbulent geophysical flows, in preparation.

Related Works and Papers

● Recent new developments

- ▶ Majda, *Introduction to turbulent dynamical systems in complex systems*, Springer, 2016.
- ▶ Majda and Qi, *New strategies for reduced-order models for predicting the statistical responses and uncertainty quantification in complex turbulent dynamical systems*, SIAM Review, 2017.

● Statistical theories

- ▶ Majda, *Statistical energy conservation principle for inhomogeneous turbulent dynamical systems*. PNAS, 2015.
- ▶ Majda and Gershgorin, *Link between statistical equilibrium fidelity and forecasting skill for complex systems with model error*. PNAS, 2011.
- ▶ Majda and Wang, *Linear response theory for statistical ensembles in complex systems with time-periodic forcing*. CMS, 2010.

● Improving imperfect model skill

- ▶ Majda and Qi, *Improving prediction skill of imperfect turbulent models through statistical response and information theory*, Journal of Nonlinear Science, 2015.
- ▶ Qi and Majda, *Low-dimensional reduced-order models for statistical response and uncertainty quantification: two-layer baroclinic turbulence*, JAS, 2016.
- ▶ Qi and Majda, *Low-dimensional reduced-order models for statistical response and uncertainty quantification: barotropic turbulence with topography*, Physica D, 2016.